SIMULATION AND COMPARISON OF MEAN FLOW, TURBULENCE AND DISPERSION IN COMPLEX TERRAIN

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- Closure schemes
- # "TRACT" experiment for the models comparison
- **#** Comparison between computed and measured data
- Analysis of the different closure schemes performances

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Conclusion and future development

INTRODUCTION

- * The integrated modelling system RMS (RAMS-MIRS-SPRAY) is used to simulate a tracer experiment carried out during the TRACT campaign (TRAsport of air Pollutants over Complex Terrain)
- RMS is based on a combination of: the meteorological model RAMS (Regional Atmospheric Modelling System), the interface code MIRS, the Lagrangian particle model SPRAY
- Herein influence of different closure schemes on turbulence fields and on dispersion is analysed



MODELLIG SYSTEM:RAMS-MIRS-SPRAY



CLOSURE SCHEMES: MELLOR and YAMADA

In RAMS

$$\frac{dE}{dt} = \frac{\partial}{\partial z} K_E \frac{\partial E}{\partial z} + P - \varepsilon$$
Level 2.5: B.L approximation, horizontal homogeneity

$$P = K_m \left[\left(\frac{\partial \overline{u}}{\partial z} \right)^2 + \left(\frac{\partial \overline{v}}{\partial z} \right)^2 \right] - \frac{g}{\theta_0} K_h \frac{\partial \theta}{\partial z} \qquad \varepsilon = \frac{(2E)^{3/2}}{\Lambda_1}$$

$$K_m = S_m l(2E)^{1/2} \qquad K_h = S_h l(2E)^{1/2} \qquad K_E = S_E l(2E)^{1/2}$$

$$(l_1, \Lambda_1, l_2, \Lambda_2) = l(A_T, B_1, A_2, B_2) \qquad l = \frac{kz}{1 + kz/L_\infty} \qquad l_\infty = 0.1 \frac{|z\sqrt{Edz}|}{|\sqrt{Edz}|}$$

$$S_m, S_h, S_E = f \left(\frac{\partial u}{\partial z}, \frac{\partial \theta}{\partial z}, E, l, A_1, A_2, B_1, B_2, C \right) \qquad (A_1, A_2, B_1, B_2, C) = (0.92, 16.6, 0.74, 10.1, 0.08)$$
From MIRS to SPRAY

$$\sigma_u^2 = (1 - 2\gamma)q^2 \qquad \sigma_v^2 = \gamma q^2 \qquad T_{Li} = \frac{K_m}{\sigma_i^2} \qquad q^2 = 2E$$

$$\overline{w^3} = -0.6 \frac{w^{s^3}}{k} \frac{z}{L} + 0.1\sigma_w^3 \qquad \text{From Chiba (1978)} \qquad \gamma = \frac{1}{3} - 2\frac{A_1}{B_1}$$

CLOSURE SCHEMES: E-I isotropic



TRACT EXPERIMENT

- Experiment performed during 7-23 September 1992
- Investigation area was about 300 x 300 km²
- He tracer was released at a height of 8 m in Sasbach in the Rhine Valley for a period of about 3 hours (from 5:02 until 7:58 of 16 September)
- **Ground network: 20 stations collected 224 g.l.c samples of 30 min**
- Heteorological measures with a sodar station located near the source
- 16 September was a sunny day with 25 C° in the valley and westerly winds



RAMS CONFIGURATION

- **4** nested grids with $\Delta x = \Delta y = 60$, 15, 5, 1.6 km, $\Delta z = 50-500$ m
- Horizontal Grid extensions 13x17, 18x18, 29x29, 62x62 points
- Stretched vertical grid extensions: 30 points up to 10000 m
- # Time step ∆t =60, 20, 6.6, 2.2 sec
- Numerical scheme: Leapfrog time differencing
- Simulation period 48 hours from 15 September 00 UTC until 17 September
 00 UTC
- Initialisation from ECMWF analysis fields
- 8 Nudging from ECMWF analysis fields every 6 hours

RAMS MIRS configuration : 4 grids locations



RAMS-MIRS simulations descriptions

4 simulations with different turbulence schemes:

-Mellor and Yamada (MY)

-E-I isotropic (EL-ISO)

-E-I for vertical direction and Smagorisky-type deformational scheme in horizontal (EL-SMA)

-Mellor and Yamada closure model with the Hanna (1982) parameterisation for wind velocity fluctuation standard deviations and Lagrangian Time (MY-HANNA)









SPRAY description

SPRAY

(Tinarelli et al., 2000, Ferrero and Anfossi, 1998) is a Lagrangian stochastic particle model for complex terrain based on three Langevin equations for the random velocities (Thomson, 1987):

$$du = a(\mathbf{x}, \mathbf{u})dt + b(\mathbf{x}, \mathbf{u})dW(t)$$
$$d\mathbf{x} = (\mathbf{U} + \mathbf{u})dt,$$

and

$$d\mathbf{x} = (\mathbf{U} + \mathbf{u})dt$$

where,

U is the mean wind velocity,

 $a(\mathbf{x},\mathbf{u})dt$

is a deterministic term depending on $P_E(x,u)$,

 $b(\mathbf{x},\mathbf{u})dW(t)$ is a stochastic term is the incremental Wiener process.

In the horizontal directions the PDF is assumed to be Gaussian. In the vertical direction the PDF is assumed to be non-Gaussian (to deal with non-uniform turbulent conditions and/or convection).

SPRAY configuration

- Bomain extension (4th RAMS-MIRS grid): horizontal 62x62 points $\Delta x = \Delta y = 1.6$ Km, vertical 30 points between 0 m and 10000 m
- # 100 particles released every 5 sec
- Ground level concentrations computed every 30 min. with cells of 500 x
 500 x 50 m³
- 2 simulations for every turbulence scheme (8 simulations): one with a constant emission period 24 hours long, one with the same emission period of the experiment

16 September: constant emission period of 24 hours



Maximum g.l.c. inside the domain







16 September 13 UTC: constant emission period of 24 hours

MY



MY-Hanna

Ground level concentrations:comparison with measured data

- # A better performance is obtained with the E-I model, both in the isotropic and non isotropic cases
- # the analysis has been performed on 7 available measurements

| Model | Mean | sigma | bias | nmse | cor | fa2 | fb | fs |
|--------------------------|------|-------|------|------|-------|-----|------|------|
| Observed | 0.6 | 1.1 | 0.0 | 0.0 | 1.0 | 1.0 | 0.0 | 0.0 |
| E-1 SMA | 0.8 | 0.7 | -0.2 | 2.0 | 0.4 | 0.1 | -0.3 | 0.5 |
| E-1 ISO | 1.2 | 1.7 | -0.6 | 2.9 | 0.6 | 0.3 | -0.6 | -0.4 |
| M&Y Hanna | 0.5 | 0.5 | 0.2 | 5.2 | -0.03 | 0.1 | 0.4 | 0.7 |
| M&Y | 1.6 | 1.7 | -1.0 | 4.1 | 0.2 | 0.0 | -0.9 | -0.5 |
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CONCLUDING REMARKS

- 38 The choice of the turbulence scheme influences both the mean field and dispersion parameters
- In unstable conditions all the models predict comparable maximum g.l.c values and the horizontal plume dispersion
- In stable conditions there are significant differences in maximum g.l.c. computed and in the horizontal plume dispersion
- Here E-I schemes show slightly better performance in forecasting measured g.l.c.