

Estimating Meteorological Inputs for Dispersion Modeling in Urban Areas

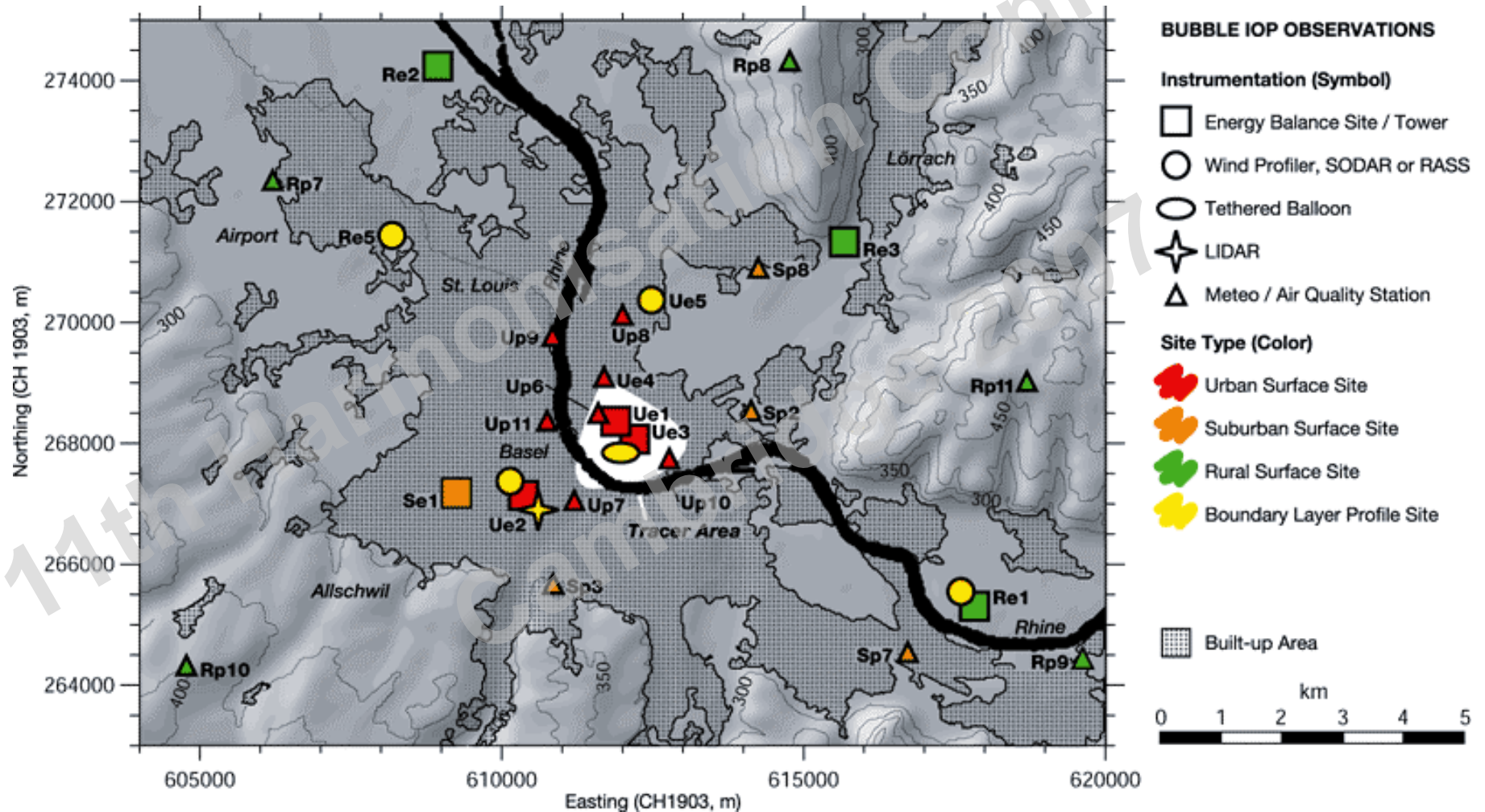
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- Meteorological models to estimate urban parameters as air flows from rural area to urban area.
 - Comprehensive TAPM-CSIRO, MM-5
 - Simple IBL models
- Measurements of mean winds and temperatures on urban towers to infer micrometeorological variables using M-O theory

TAPM-The Air Pollution Model

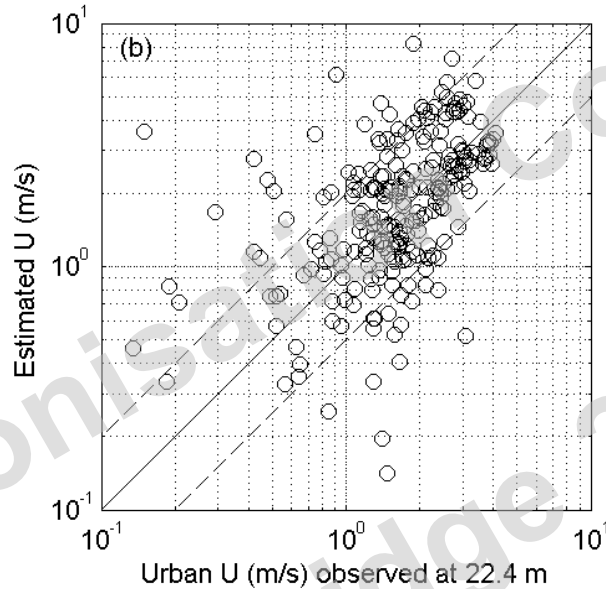
- TAPM used to simulate July 10-11, 2002 during the BUBBLE experiment with fine resolution of 0.5 km.
- The main urban measurements tower, Basel-Sperrstrasse, 32 m high, located inside a street canyon in area with homogeneous, residential building blocks, with mean height of 14.6 m AGL
- The rural site, Village Neuf, located about 6.5 km NNW of the urban site, measured flow and turbulence at 3.3 m AGL over bare soil in an agricultural area ($= 0.07$ m).

BUBBLE Experiment

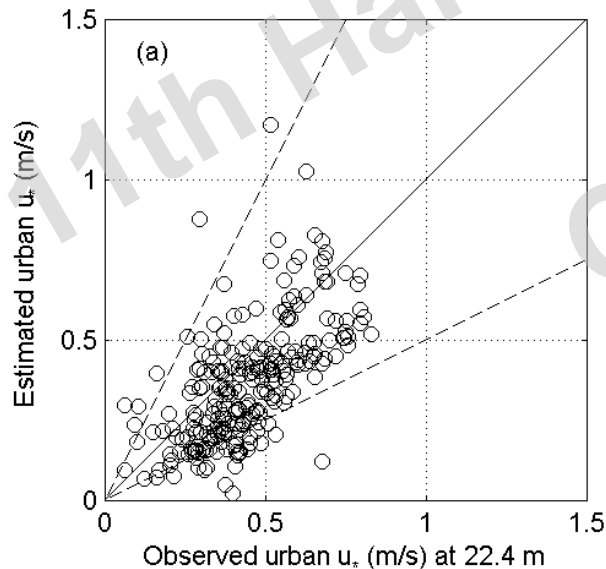


Results from TAPM

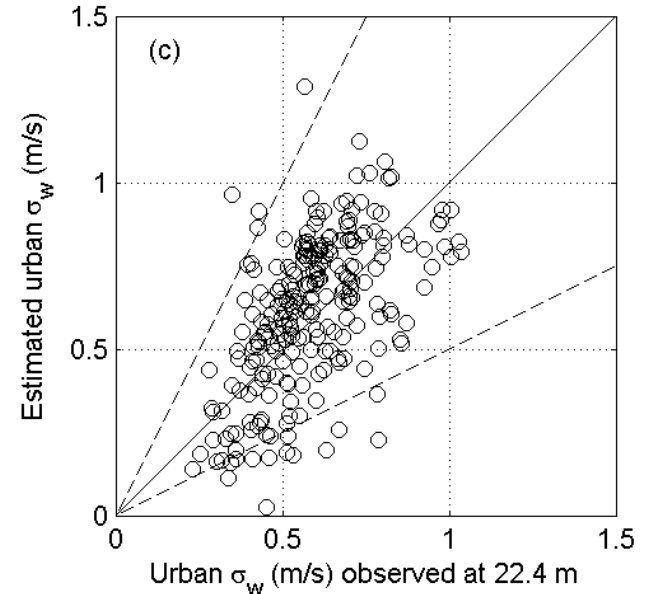
Mean Wind



Friction velocity

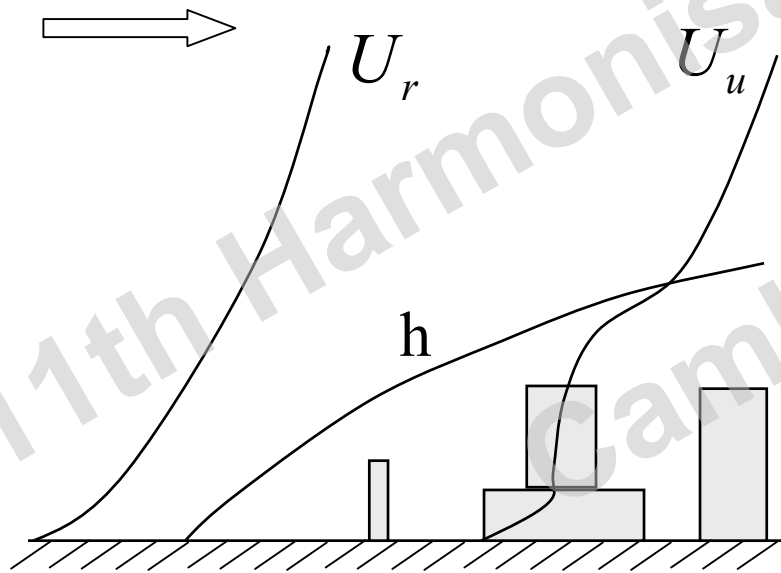


Vertical velocity fluctuations



Internal boundary layer model

Internal boundary layer model



$$\frac{dh}{dx} = \frac{\sigma_w}{U} \varphi \left(\frac{h-d}{L_u}, z_{ou} \right)$$

$$L_u = \infty \text{ when } L_r > 0$$

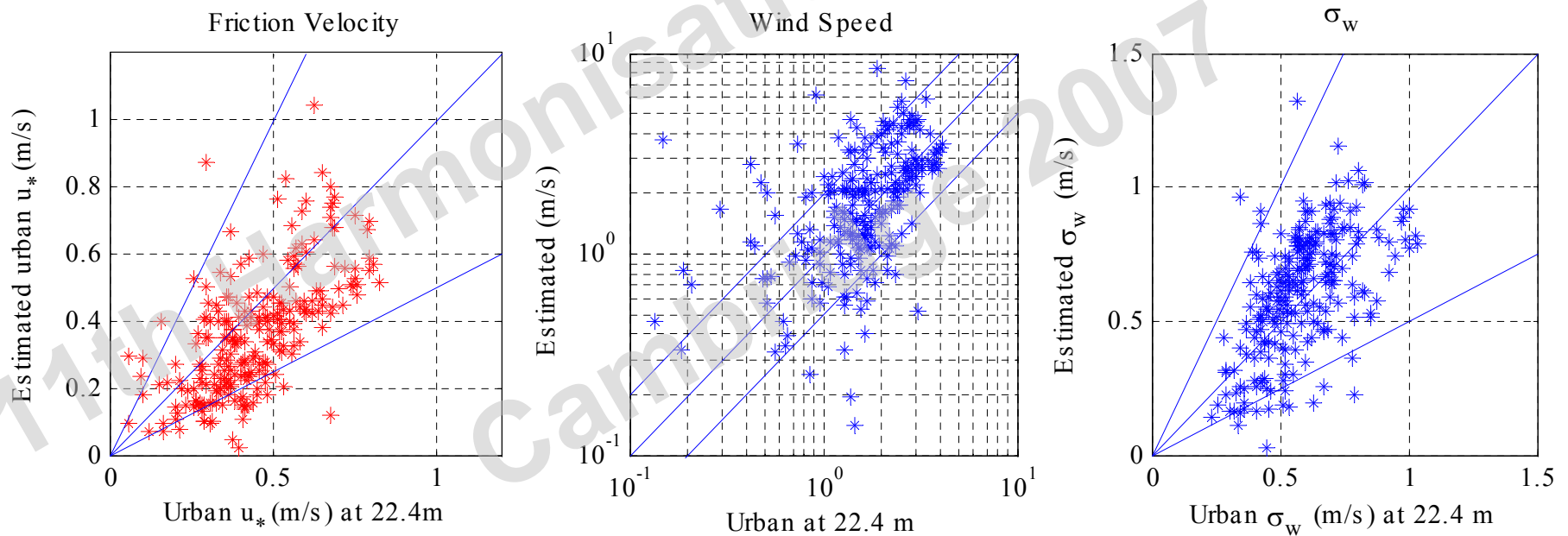
$$L_u = L_r \text{ when } L_r < 0$$

$$U_u(h) = U_r(h)$$

$$u_{*u} = \frac{kU_u(h)}{\ln \left(\frac{h-d}{z_{ou}} \right) + \varphi_m \left(\frac{h-d}{L}, z_{ou} \right)}$$

Evaluating Urban Micrometeorological Variables with IBL model

Measurements at BSPR from BUBBLE



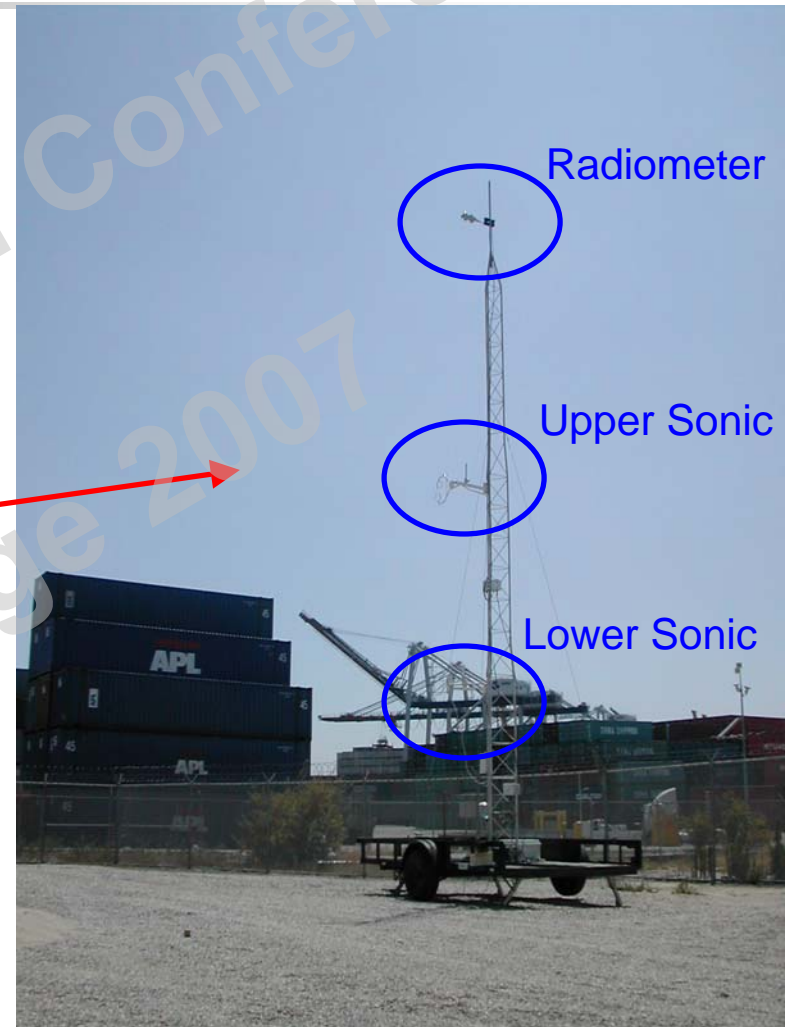
Remarks on using Met Models

- Can provide estimates of u_* , and σ_w
- Heat flux and boundary layer heights are unreliable
- Need tuning with moisture parameter and other surface parameters

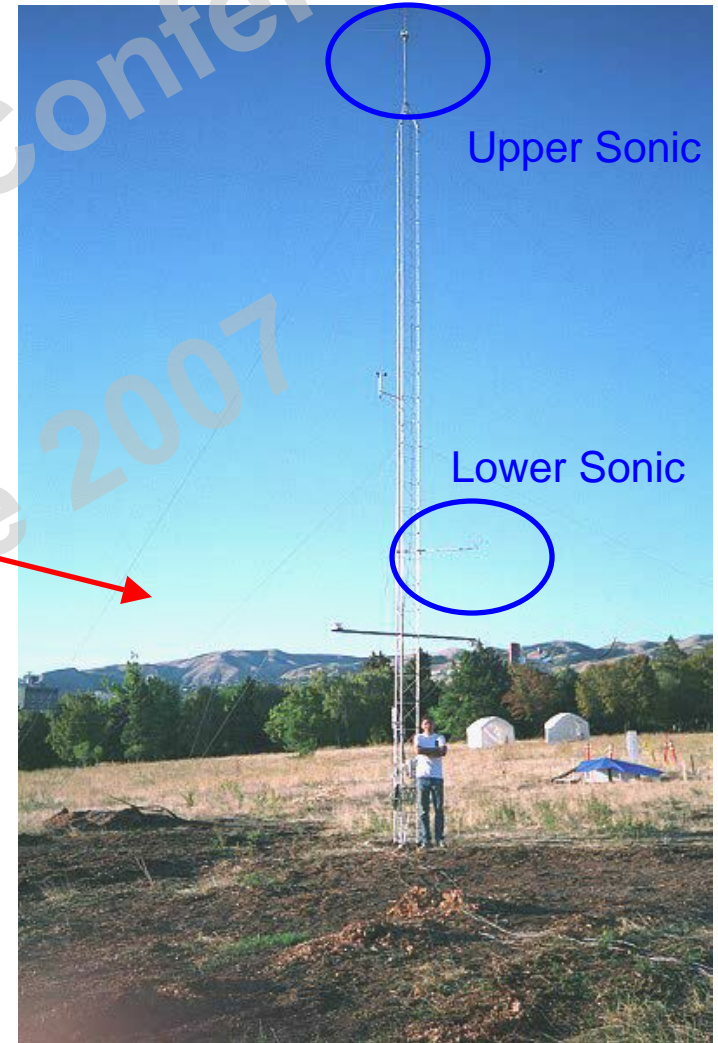
Using Urban Measurements

- Make measurements of mean winds, temperature, temperature fluctuations, at one or two levels on a tower
- Use M-O similarity to estimate surface friction velocity, M-O length, and standard deviations of turbulent velocities

Wilmington Instrumentation



VTMX Instrumentation



Stable Conditions

One level of wind: Assume θ_* is constant (Venkatram, 1980)

$$u_* = C_D u(z_r) \left\{ \frac{1}{2} + \frac{1}{2} \left[1 - \left(\frac{2u_0}{C_D^{1/2} u} \right)^2 \right]^{1/2} \right\}$$

$$\sigma_w = 1.6u_*$$

and

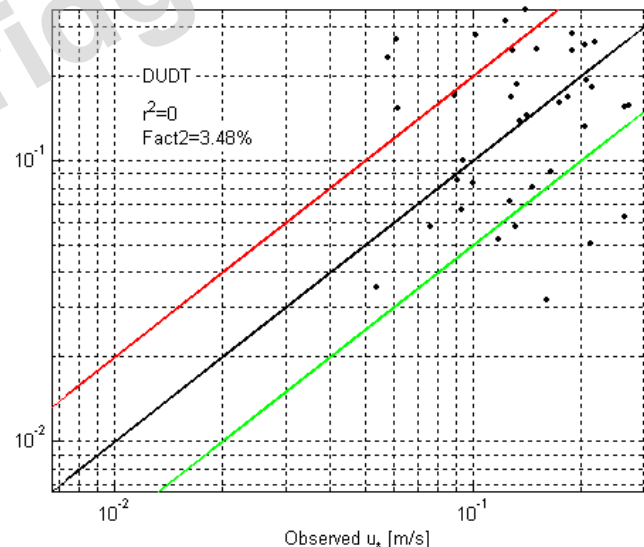
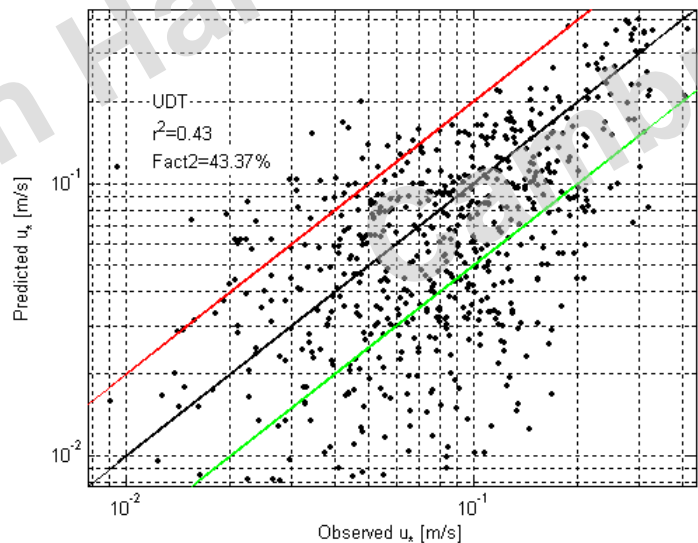
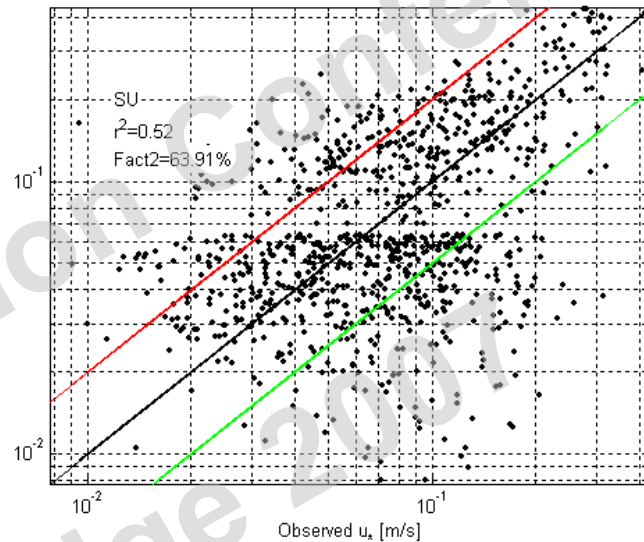
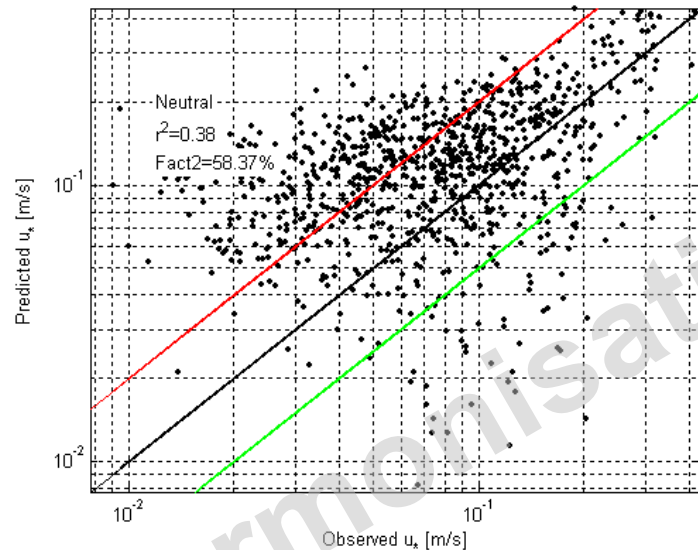
$$\sigma_v = 1.9u_*$$

$$\theta_* = 0.5\sigma_T$$

Two levels of temperature: Bulk Richardson number (Irwin and Binkowski, 1980)

Results-Stable Conditions

Site: Wilmington, CA (Los Angeles Metropolitan Area)



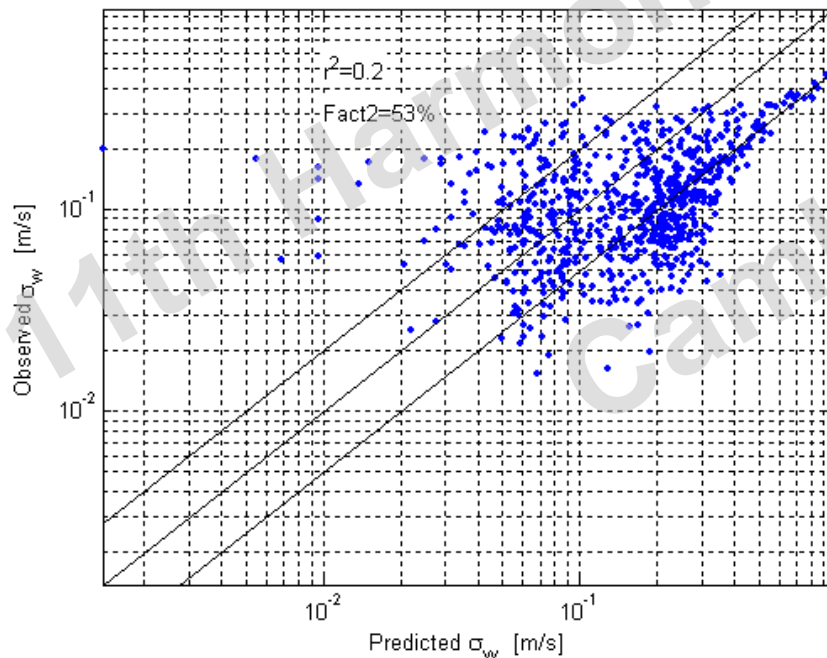
Results-Stable Conditions

$$u_* = C_D u(z_r) \left\{ \frac{1}{2} + \frac{1}{2} \left[1 - \left(\frac{2u_0}{C_D^{1/2} u} \right)^2 \right]^{1/2} \right\}$$

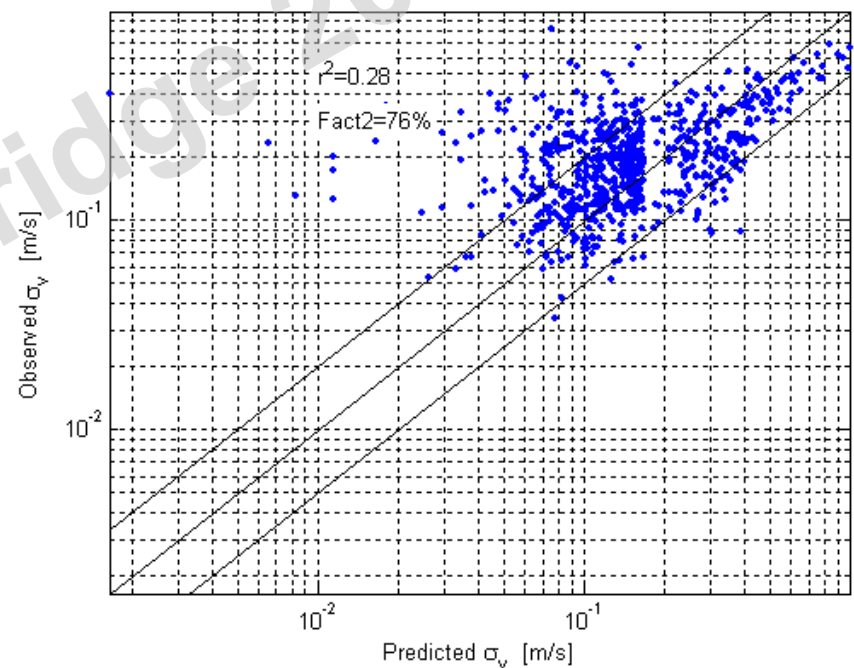
$$\sigma_w = 1.6u_*$$

$$\sigma_v = 1.9u_*$$

Wilmington: SU Observed vs. predicted σ_w on u_* with θ_s on θ_0 - Stable



Wilmington: SU Observed vs. predicted σ_v with θ_s constant - Stable



Unstable Conditions

$$Q_0 = \left(\frac{\sigma_T}{0.95} \right)^{3/2} \left(\frac{gkz_r}{T_0} \right)^{1/2}$$

$$\sigma_w = \left(\sigma_{ws}^3 + \sigma_{wc}^3 \right)^{1/3}$$

$$\sigma_{ws} = 1.3u_*$$

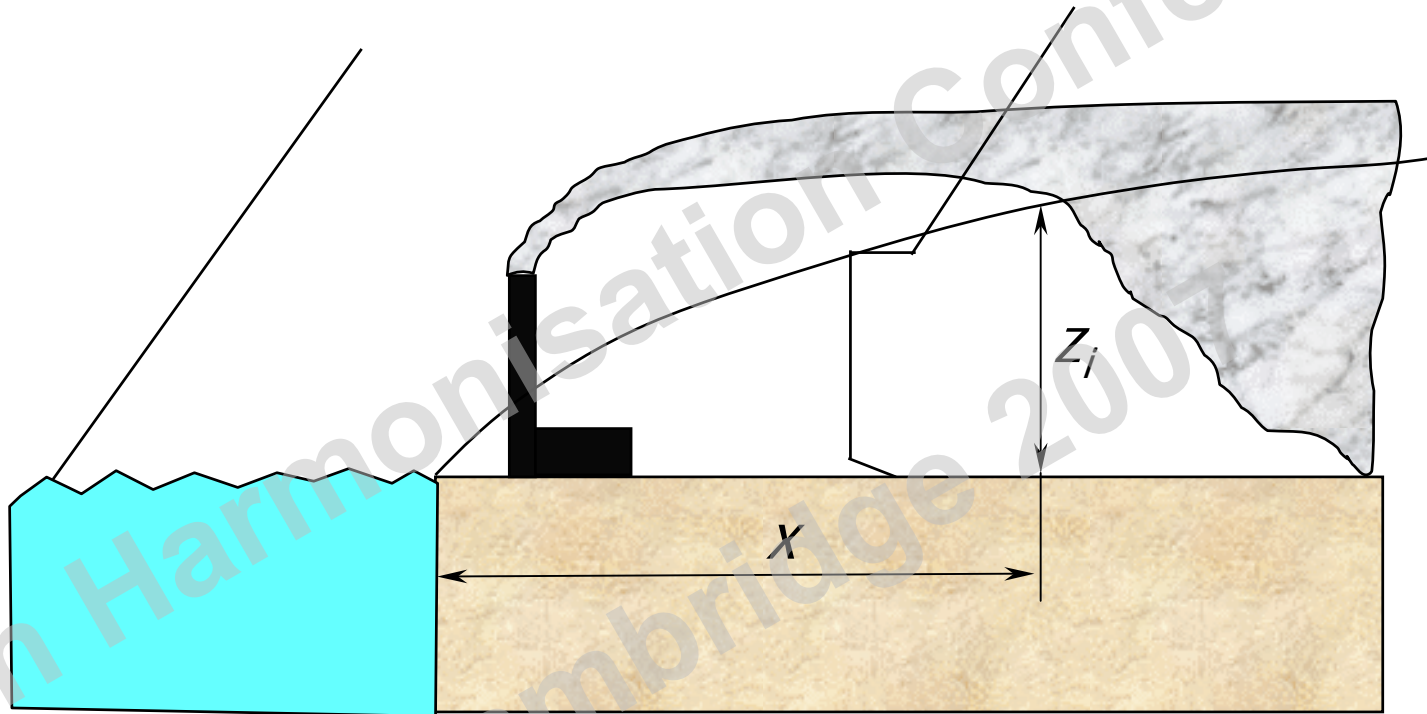
$$\sigma_{wc} = 1.3 \left(\frac{g}{T_0} Q_0 z_r \right)^{1/3} \quad \text{for } z_r \leq 0.1z_i$$

$$= 0.6w_* \quad \text{for } z_r \geq 0.1z_i$$

where the convective velocity scale w_* is

$$w_* = \left(\frac{g}{T_0} Q_0 z_i \right)^{1/3}$$

Wilmington-Shoreline Meteorology



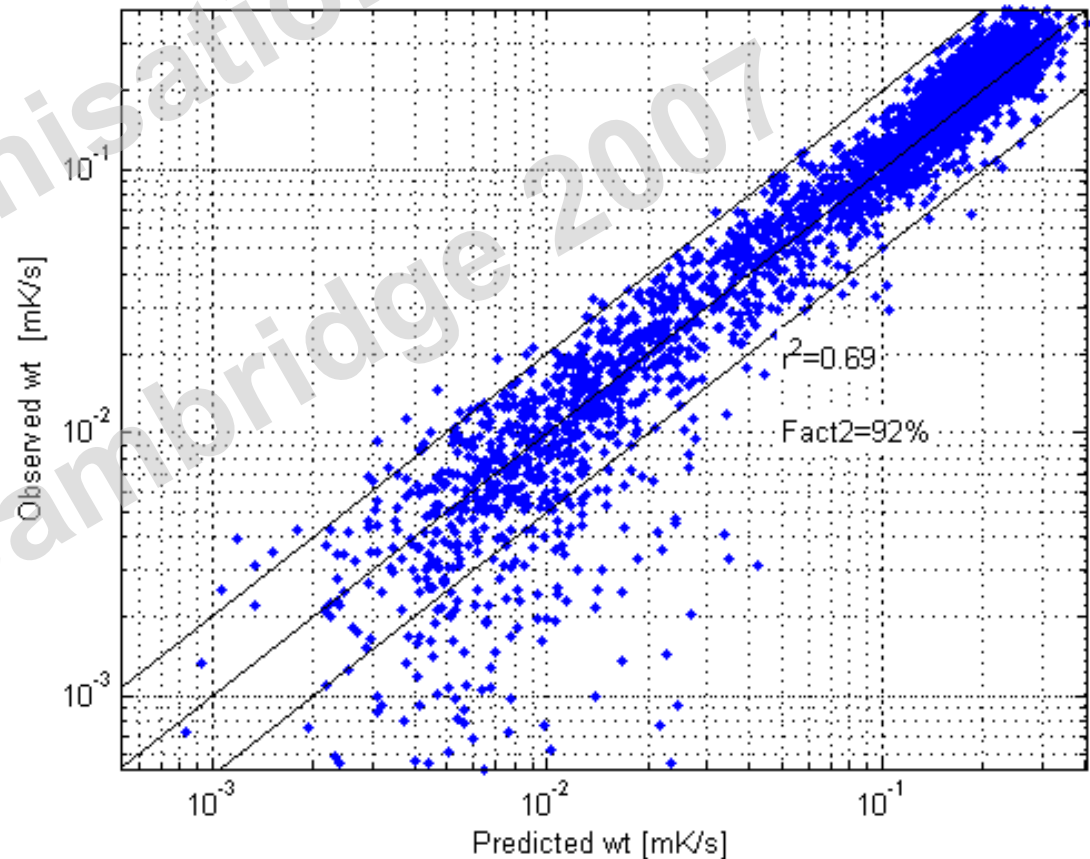
$$z_i = a \left(\frac{Q_0 (x + x_0)}{U\gamma} \right)^{1/2}$$

$$w_* = Q_0^{1/2} \left(\frac{g}{T_0} \right)^{1/3} \left(\frac{\alpha(x + x_0)}{U\gamma} \right)^{1/6}$$

Results-Unstable Conditions

$$Q_0 = \left(\frac{\sigma_T}{0.95} \right)^{3/2} \left(\frac{gKZ_r}{T_0} \right)^{1/2}$$

Wilmington: Observed vs. predicted wt on σ_θ - Convective



Results-Unstable Conditions

$$u_* = \kappa u \frac{1 + d_1 \ln(1 + d_2 d_3)}{\ln(1/r_h)}$$

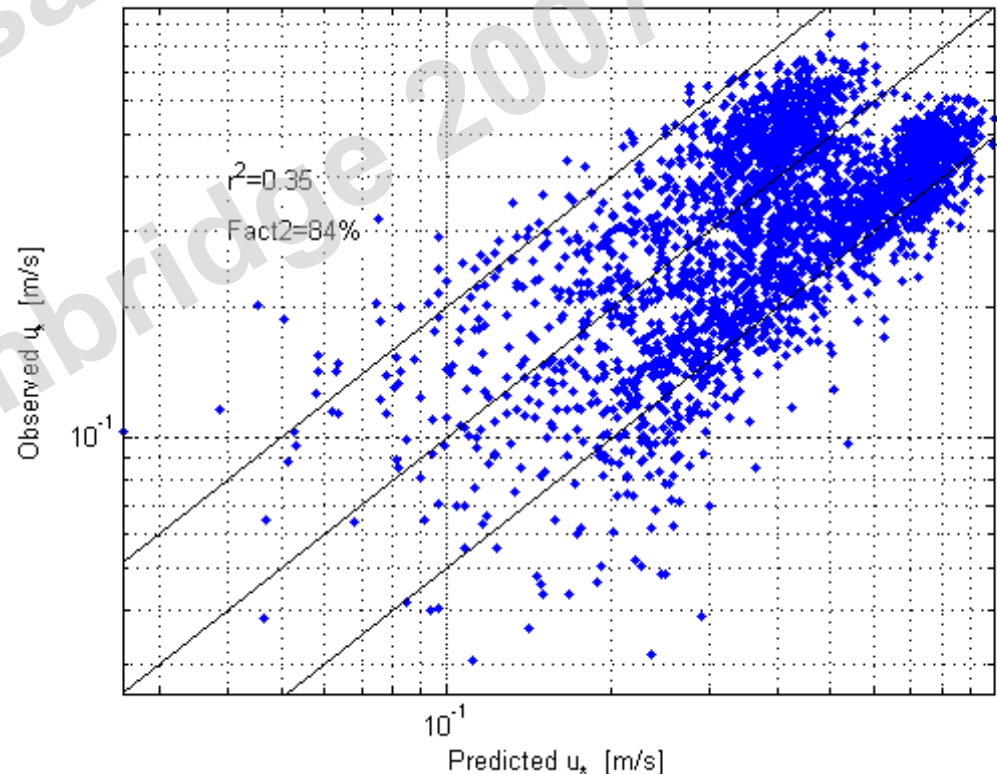
$$r_h = \frac{z_0}{z_r - d}$$

$$d_1 = \begin{cases} 0.128 + 0.005 \ln(r_h) & \text{for } r_h \leq 0 \\ 0.107 & \text{otherwise} \end{cases}$$

$$d_2 = 1.95 + 32.6 r_h^{0.45}$$

$$d_3 = \frac{Q_0 \kappa g (z_r - d)}{T_0}$$

Wilmington: Observed vs. predicted u_* on σ_θ - Convective



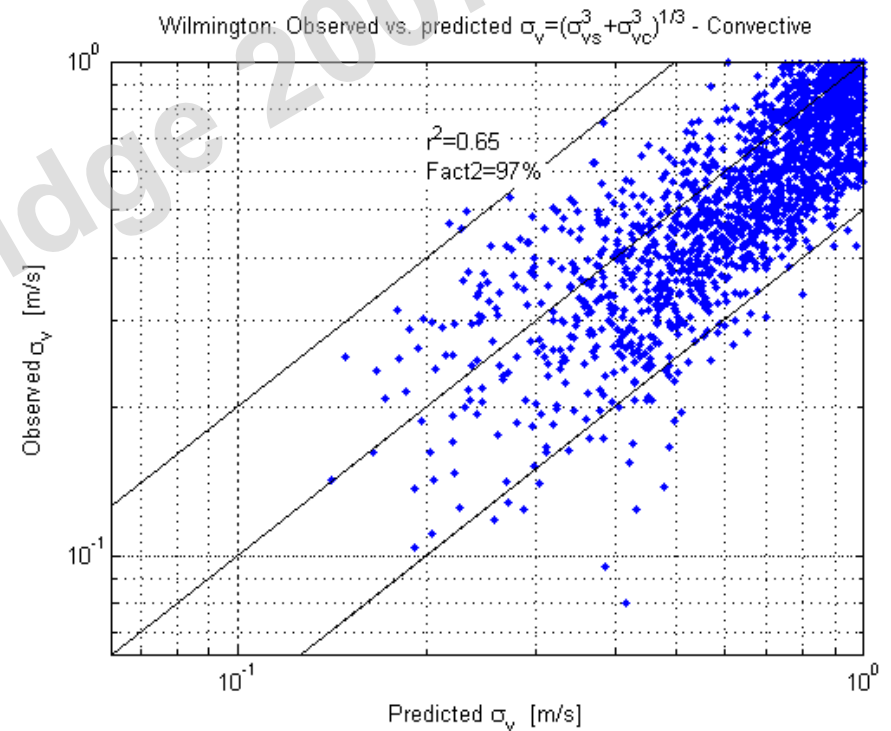
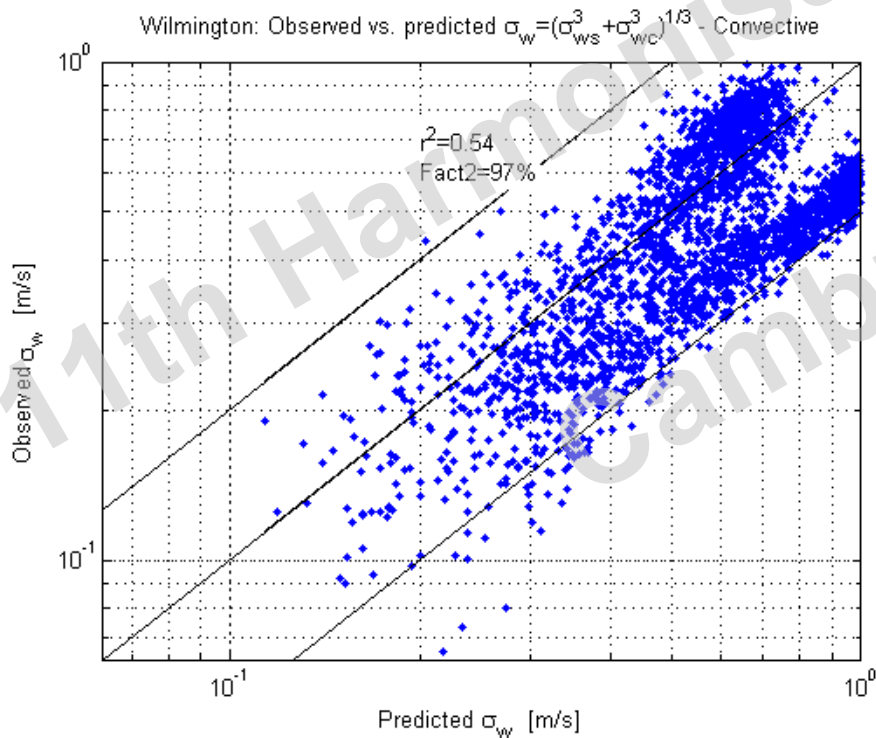
Results-Unstable Conditions

$$\sigma_{ws} = 1.3u_*$$

$$\sigma_{wc} = 1.3 \left(\frac{g}{T_0} Q_0 z_r \right)^{1/3}$$

$$\sigma_{vs} = 1.9u_*$$

$$\sigma_{vc} = 0.6w_*$$



Conclusions

- ◆ Comprehensive and IBL models provide reasonable estimates of urban parameters. However, calculation of heat flux is uncertain.
 - ◆ Moisture parameter dominates heat flux calculation
- ◆ One or two level measurements of winds and σ_T can be used to estimate urban parameters. More testing is required.
 - ◆ Two levels can increase errors

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